

APPENDIX 6.2

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THE FORAMINIFERAL SEQUENCE IN VOLADOR-1

BY DAVID TAYLOR

FOR: SHELL DEVELOPMENT (AUSTRALIA) PTY. LTD.

R 4688

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<u>VOLADOR # 1</u>.

SUMMARY OF FORAMINIFERAL SEQUENCE

Depth (m)	ZONE*	AGE*	PALEOENVIRONMENT	E-Lo Pick
1325	A-4	Early Pliocene	Shelf Edge (≃250m)	
1475	<u></u>		Upper Slope	
to	to		carbonates (400-250m)	
1785	С	Mid Miocene		
1925			Slumping of carbonate	
to	D-1	Mid Miocene	down slope with	
2555			canyon fills	
·····	····· (2 m.	v.)······	······	w256
2564	E-1	Mid to	Carbonates at base	
to	to	Farly	of continental	
2650	E-2	Miccene	$slope(\approx 1000m)$	
2050	£-2	Miocene		
2671			Deep oceanic basin	200
2071	r	Early	errhenete with sporadia	
	τn		Califonale with sporadic	
to	-	Miocene	annuanian Jua ta	
to 2828	G	Miocene	corrosion due to	
to 2828 	G J-2	Miocene 4 m.y.) Early Oligopopo	corrosion due to proximity of C.C.D. (2000-1000m). WWWWWWWWWWWWWWW Carbonates from rapid transgression & eustatic	vv284
to 2828 2850 to 2935	G J-2	Miocene 4 m.y.) Early Oligocene	corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m).	~284 293
to 2828 2850 to 2935 2941	G J-2 J/K	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m)	₩284 293
to 2828 2850 to 2935 2941	G J-2 J/K	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) 2 2 2</pre>	~293 293 2294
to 2828 2850 to 2935 2941	G J-2 J/K	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ????	₩284 293 ?294
to 2828 2850 to 2935 2941 2941 2949	G J-2 J/K ? No	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ???? Barrier barred estuary/ laconal system: apoxic</pre>	~293 293 ?294
to 2828 2850 to 2935 2941 2941 2941 2949 to 2010	G J-2 J/K ? No planktoni	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ???? Barrier barred estuary/ lagoonal system; anoxic for polybaling (<10m)</pre>	₩284 293 ?294
to 2828 2850 to 2935 2941 2941 2949 to 3010	G J-2 J/K ? No planktoni	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ???? Barrier barred estuary/ lagoonal system; anoxic & polyhaline (<10m).</pre>	~293 293 ?294
to 2828 2850 to 2935 2941 2941 2941 2949 to 3010 ?	G J-2 J/K ? - No planktoni ? -	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary "Early Tertiary" cs. arenaceous forams. 	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ???? Barrier barred estuary/ lagoonal system; anoxic & polyhaline (<10m). ?????</pre>	~293 293 ?294 ?
to 2828 2850 to 2935 2941 2941 2941 to 3010 2947	G J-2 J/K ? No planktoni ? ?	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary "Early Tertiary" cs arenaceous forams. ? ? ? ?	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ? ? ? Barrier barred estuary/ lagoonal system; anoxic & polyhaline (<10m). ? ? ? ? ?</pre>	~293 294 ?
to 2828 2850 to 2935 2941 2941 2941 2949 to 3010 ? 2947 to	G J-2 J/K ? No planktoni ? No microf	Miocene 4 m.y.) Early Oligocene Eo/Oligocene Boundary 	<pre>corrosion due to proximity of C.C.D. (2000-1000m). Carbonates from rapid transgression & eustatic sea level rise onto shelf platform (40m-<200m). Commencement of early Oligocene Transgression coincidental with tectonic adjustment (<40m) ???? Barrier barred estuary/ lagoonal system; anoxic & polyhaline (<10m). ? ?</pre>	~293 293 ?294 ?

- foraminiferal distributions on Table 2, with reliability of zonal boundary picks given on Table 1.
- Interpretations based on benthonic foraminifera (see Table 3); total microfauna and other sediment grains (see Table 4). Paleodepth estimates given in parentheses.

MAX(2 m.y.) For Hiatus with time span in parentheses.

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BIOSTRATIGRAPHY.

This discussion is based on the distribution of Tertiary planktonic foraminifera in the Gippsland sector of the Tasman Sea. Provence. The biostratigraphic letter scheme is in accordance with those for other sectors of the Provence, as outlined by Jenkins (1974), and Srinivasan & Kennett (1981).

EARLY TERTIARY - 3010 to 2949: No planktonic foraminifera found, but the arenaceous benthonic foraminiferal assemblages were similar ' to arenaceous assemblages, which were common in early Tertiary sediments in all three Bass Strait Basins (Taylor, 1965).

EO/OLIGOCENE BOUNDARY - 2041m: Only planktonic foraminifera present were two specimens of *Globigerina angiporoides angiporoides* which ranged from latest Eocene into the basal Oligocene.

<u>BASAL OLIGOCENE - ZONE J-2 - 2935 to 2850m</u>: Contains planktonic assemblages typical of the basal Oligocene (Zone J-2) of the Tasman Sea region; especially *Globigerina brevis* and *Globorotalia gemma* (? syn. or \equiv *G*. postcretacea).

EARLY MIOCENE - ZONES G at 2828m and the OLIGO/MIOCENE HIATUS at 2840m (E-log): At 2828m, the presence of Globigerinoides trilobus with the Tasman Sea early/mid Miocene Globorotalia suite of G. miozea miozea, G. praescitula and G. zealandica indicates a biostratigraphic position some distance above the base of the Miocene; that is Zone G, rather than the basal Miocene Zone H-1. Such a biostratigraphic placement of the sidewall core at 2828m automatically implies a hiatus existed between the early Oligocene at 2850m and the early Miocene. The E-log pick for this event in Volador # 1 is at 2840m. This event was an ubiquitous one in deep water Oligo/ Miocene sequences in the Gippsland Basin as well as in the whole Tasman Sea Provence (refer Loutit & Kennett, 1981a). In Volador, the time span of the Hiatus was some 12 to 14 million years in length, compared with a maxima of 20 million years recorded in Hapuku # 1. However, in marginal, shallow water, Gippsland sequences, sedimentation was accumulating during this interval (e.g. at Lakes Entrance - refer Jenkins, 1974, Table 3). The causes of this hiatus will be discussed in the Paleoenvironment section of this report.

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EARLY/MID MIOCENE - ZONES G,F.E-2 & E-1 - 2828m to 2564m: A deep water sequence of early Miocene carbonate extends from Zone G, through Zone F (with Globigerinoides bisphericus \equiv G. sicanus), to Zone E-2 at the top of the early Miocene with members of the Praeorbulina glomerosa Group. Sediment accumulation was also evident at the very base of the mid Miocene (Zone E-1 with Orbulina suturalis); before the sequence was abruptly truncated by mid Miocene slumping and submarine canyon cutting.

MID MIOCENE - BASE of CANYON FILL SEQUENCE - ZONE D-1 at 2555m: Probably all the mid Miocene Zone D-2 sediment was removed by slope slumping and/or the initiation of canyon cutting. A hiatus, with a time span of some 2 million years was evident in the Volador sequence, with an E-log pick at 2563m.

MID MIOCENE - ZONE D-1 - CANYON FILL SEQUENCE 2555 to 1925m: A thick sequence of Zone D-1 assemblages, which fluctuate in the quality of preservation. The degree of diagenesis was proportional to the ability to specifically identify specimens and thus positively identify a Zone D-1 assemblage in each sample. This is demonstrated by the tabulation below and in several columns on Table 4.

Depth (m)	Sidewall Core	ZONE	Reliability Rating (refer Table 1) & Preservation Quality				
1925	# 43	indeterminate		poor			
2001	# 42	D-1	(1)	moderate			
2144	# 40	indeterminate		poor			
2300	# 37	indeterminate		poor			
2350	# 36	D-1	(2)	poor			
2426	# 34	D-1	(1)	moderate			
2452	# 33	indeterminate		poor			
2500	# 31	D-1	(0)	moderate			
2525	# 30	D-1	(0)	moderate			
2555	# 29	D-1	(1)	moderate			

Samples with a high confidence rating of zero (O), contained the *Globorotalia* peripheroacuta morphotype of the *G.foshi* Group; whilst those with a moderate rating of one (1) contained members of the Tasman Sea Provence *Globorotalia* suite, including the highest appearance of *G. conica* at 1925m.

MID to LATE MIOCENE to EARLY PLICCENE - ZONES C.B-2 and A-4 - 1785 to 1325m:

Because of poor sidewall core recovery, the upper part of the Gippsland sequence was poorly represented in Volador # 1. Zone B-1 was not designated but was probably present in the unsampled interval between 1475 and 1325m. The sample at 1325m contained an almost complete list of species for a Zone A-4 planktonic assemblage, bearing witness of the early Pliocene transgression onto the southern Australian Margin.

PALEOENVIRONMENT.

These interpretations are based on the analysis of the benthonic foraminiferal faunas, as well as all faunal elements (including planktonic foraminifera) and other sediment grains in the prepared residues (size >.075mm). Paleodepths and other physico-chemical parameters are deduced from the benthonic assemblage by a combination of both comparison with faunal distribution in sediments of similar age and by using present day distribution of species analogues.

3010 to 2949m "EARLY TERTIARY" ESTUARINE: These arenaceous benthonic foraminiferal assemblages were, as elsewhere in the Bass Strait region, indicative of poorly oxygenated water with severe salinity fluctuations (Taylor, 1965). Distinctly biogenic pyrite and indications of sediment bioturbation were associated with these faunas, as were pyritized discs and spheres which could be attributed to diatoms. Such fossil features are characteristics of the carbon-rich, mudstones and micaceous quartz sandy siltstones of the Flounder Formation of the Gippsland, the Demons Bluff Formation of the Bass Basin and the Johanna River Sand and Dilwyn Formation of the Otway Basin. The facies similarities suggest the development of a series of barrier-barred estuarine/lagoonal systems. However, the facies was extremely diachronous, ranging in age from latest Paleocene to Eocene across the three Bass Strait Basins.

By this comparative reasoning, I regard the Flounder Formation to fill a series of estuarine scours rather than deeply incised channels.

<u>2941 to 2850m - EO/OLIGOCENE - RAPID TRANSGRESSION</u>: The four assemblages within this interval demonstrate rapid environmental change; from the

initiation of the marine transgression onto an exposed surface at 2941m to total inundation by shelfal seas and establishment of a biogenic carbonate, sedimentary regime. Both the planktonic and benthonic foraminiferal assemblages indicate that this transgression was isochronous with and resulted in similar environmental responses as the Whaingaroan Stage in New Zealand (Loutit & Kennett, 1981b).

The "Vail Coastal On-lap Curve" (e.g. Loutit & Kennett, l.c., p.1596) expresses this Whaingaroan transgressive event in terms of eustatic sea level rise. However, the increase in paleo-water depth in the Gippsland Basin Deep was too exaggerated in vertical scale, to have resulted purely from eustatic sea level rise. Vulcanism and tectonic uplift occurred in the East Gippsland Highlands in latest Eocene Times (Wellman, 1974), so that compensatory subsidence could explain the rapidity of basin deepening (refer below for further discussion).

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OLIGOCENE/EARLY MIOCENE HIATUS - 2840m (E-log): The sidewall cores above this hiatus contained an extremely deep water benthonic foraminiferal assemblage associated with Zones G & F (early Miocene) planktonic faunas. The benthonic faunas were often dominated by morphologically primitive agglutinated foraminifera; especially the branching tubed form Rhabdammina These benthonic assemblages closely resemble those recorded abyssorum. in sediments of the same age from DSDP Site 206 in the Tasman Sea by Hayward & Buzas (1979). By analogy, these Volador faunas are typical of modern assemblages inhabiting the continental rise between 2000-4000m. This depth is confirmed by comparative methods, in that the interpreted depths are "consistent with inferred little change in water depth (3000m) at DSDP 206 between the early Miocene and now" (Hayward & Buzas, l.c., p.24). A higher depth estimation of between 2000 and 1000m is given for the Volador # 1 early Miocene as faunal modification may have occurred due to upwelling of carbonate deplete water, thus raising the C.C.D. at the base of the continental rise.

Effects of the corrosion due to differential dissolution of CaCO₃ are noted in the Volador early Miocene with the degree of corrosion fluctuating sporadically (see Table 4). Thus carbonate sedimentation was in the 'proximity of the C.C.D. throughout the early Miocene. However, Loutit & Kennett (1981a, fig. 1-4) show that the C.C.D. was deeper during the late Oligocene to early Miocene than it was during the early Oligocene and Eocene, or during the mid to late Miocene. Loutit & Kennett (1981a & b and Kennett references therein) account for the hiatus in terms of paleoceanographic changes associated with the opening of the Southern Ocean, causing northern movement of high velocity deep currents to penetrate into the Tasman Sea. The non-accumulation of sediment was thus explained for in terms of non-deposition or ertosion. But the time span of the hiatus is by no means uniform throughout the Gippsland Basin. This inconsistency is even apparent in wells within the same structure. Instances of the time span variation are:-

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Zero along Basin Margins (e.g. Lakes Entrance) 8-10 million years in Kingfish, Fortesque, Cobia Wells 12-14 million years in Volador # 1 17 million years in Flounder # 1 and 20 million years in Hapuku # 1.

This data supports the contention that upwelling of the C.C.D. was responsible for the absence of some if not all Oligo/Miocene carbonate sedimentation. The extent of the upwelling would have been a function of the submarine topography in the proximity of any one site.

2828 to 2564m - EARLY to MID MIOCENE BASIN DEEP CARBONATES: The deep water environmental interpretation for the Zone G & F interval (2828 to 2671m) is discussed above. Gradual shallowing of the depositional surface occurred and is most marked at and above 2650m, during the early/mid Miocene transition of Zones E-2 and E-1. A situation at the base of the continental slope (\approx 1000m) is envisaged.

2555 to 1925m - MID MIOCENE CANYON FILL CARBONATES: High energy sedimentation is evident at and above 2555m, compared with generally low energy regimes of the continental rise below 2564m. Recycling of shallow shelf benthonics and size and shape sorting of planktonics are features of assemblages at and above 2555m. This 2555m sample was just above the truncation of the biostratigraphic sequence during the mid Miocene (at 2563m on E-log). All these features are typical of submarine cutting and subsequent carbonate filling in Gippsland offshore sequences. The time span for this Zone D-l fill with a thickness of at least 600m was less than 1.5 million years. This accumulation rate was remarkably high. No doubt much of the accumulation was slumped material from an unstable shelf edge and upper slope. REFERENCES.

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Table 1

MICROPALEONTOLOGICAL DATA SHEET

ВΑ	S I	N:G	IPPSLAND				ELEVA	ATION: KB	: _+2	25.3 GL:	-26	0				
WEL	l na	ME:V	DLADOR # 1				TOTAL	DEPTH:								
			HIG	ΗE	ST D	АТ	A	LOWEST DATA								
A G E ZONULES		Preferred Depth	Rtg	Alternate Depth	Rtg	Two Way Time	Preferred Depth	Rtg	Alternate Depth	Rtg	Twc T					
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		^A 2														
ι υ		A ₃					'		ļ							
PLI(CENI		^A 4	1325	0				1325	0							
F	ធ្ន	^B 1														
	LAT	^B 2	1475	2				1475	2							
		<u> </u>	1700	0				1785	1							
ш	ы Г	^D 1	2001	1				2555	1	2525	0					
z w	۵	2										• • • • • • • • • • • • • • • • • • • •				
U U	а н	E ₁	2570	0				2570	0							
0	Σ	E ₂	2590	0				2650	0							
W		F	2671	1				2750	1							
	REY	G	2760	1				2828	0							
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No.	۲ د	^I 2														
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0	EAR	^J 2	2850	1				2941	2	2935	0					
Х Е	2	ĸ	2941	2				2941	2							
	i	Pre-K	2947	2				3010	2							
CON	лмen	TS Pre-K	interval d	oes	not conta:	Ln p	lanktoni	c foramin	ifer	al; but						

arenaceous fauna & diatoms were present and together with litho-

facies suggest Flounder Formation.

N.B. Hiatus between J-2 & G (time span =12-14m.y.) and between

E-1 & D-1 (time span ~2 m.y.). Absence of B-1 was probably

an artifact of sampling.

CONFIDENCE RATING:	0: 1· 2: 3· 4·	SWC or Core SWC or Core SWC or Core Cuttings Cuttings	 Complete assemblage (very high confidence). Almost complete assemblage (high confidence). Close to zonule change but able to interpret (low confidence). Complete assemblage (low confidence). Incomplete assemblage, next to uninterpretable or SWC with death suspicion (very low confidence).
			depth suspicion (very low confidence).

NOTE

If an entry is given a 3 or 4 confidence rating, an alternative depth with a better confidence rating should be entered, if possible. If a sample cannot be assigned to one particular zone, then no entry should be made, unless a range of zones is given where the highest possible limit will appear in one zone and the lowest possible limit in another.

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ENVIRON-	Barred	Transgressive	CCEANIC BA	STN DESD	PACE of PT	E and Chong	CUELD
MENT	/lagoon	shelf platform			BASE OF RIS	SE and SLOPE	SHELF
SIDEWALL CORES Depth in metres	Armodiscus parri Bathysiphon angleseaensis Naplophraymoides incisa H. paupera H. paupera Diaca f anharen 7 pinome	Bollvina finlayi B. anastomosa Bulladna truncana Plectofrondicularia whaingaroica Floctofrondicularia whaingaroica Floctularia "trochus" Bathysiphon spiporcelaineous) Rhabdammina abyssorum	Ammodacus incertus Signodulamina perifica Globobulamina perifica Gyroddina zealandica Oridorgalis tenera Vulvulina pennetula Pistominelia exigua Cyclammina compressa	Aumoglobigerina sp. Hyperannina subbodosum Eggerella bradyi Textularia goesii Cibicides mundulus C. wuellerstorfi C. wuellerstorfi Bibicides karreriformis Hyperammina cylindrica	Recurvoides app. Melonis barleeanus Melonis barleeanus Fisaurina app Lagena app Nopkinsina mioindex Cibicides temperatus Rhabdammina algeeformis Bullino marinara	Saccammina sp. Lenticulina sp. Lenticulina spp. Oridorsalis umbonatus Cassidulina subglobosa Cassidulina levigata Euuvigerina bradyi Trifarina bradyi Sippouvigerina proboscidae	Cibicides vortex Cibicides vortex Anomalina macroglabra Pseudoclavulina rudis Textularia carinata Bollvinita guadrilatera Virgulina spp. 2 y 5
1325.0 ₊ 1475.0 ₊		×				* x * * * * * * * * * * * * * * * * * *	x •••• /
1700.0. 1785.0. 1925.0. 2001.0.	indet				• •	× ×	
2144.0 ₊ 2300.0 ₊ 2350.0 ₊	indet indet					x	
2426.0.	indet	•	•	•	×	* * * * * *	
2525.0. 2555.0.		•	• x	×	• • •	x · · · x	RRR
2564.0 ₊ 2570.0 ₊	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	•••••••••		~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	****	~~~~~	E
2580.0. 2590.0.			•	×	D x * x		
2610.0.		د	•	x x	• • _x		E
2631.0. 2640.0.		x	• • • •	* x • • • • • •	× • •		
2659.0. 2671.0.			•	• 			
2690.0. 2690.0.		x • D x •	· · ·	• x x x			
2710.0		ם ס	* • • • •	x ****			
2740.0.		- x D	•	•			
2760.0		X	* x * * *				
2790.0+ 2800.0+		:	•				
2810.0+ 2820.0+ 2825.0+		× D	• •				
2828.0.	R R	, mmm	••	······	······		·····
2850.0 ₊ 2890.0 ₊ 2935.0 ₊		• • • •					J
2941.0+		indet7	7	- 7 7		, <u> </u>	
2970.0, 2990.0, 3000.0, 3010.0, 3019.3, 3019.5, 3030.0,	x • • • • • • • • • • • • • • • • • • •						
<u>KEY</u> : • - D - 7 -	<20 specime >20 specime Dominant >6 determinati	ns ns 0% specimens on queried	R = Recycled specie N.F.F. = no forami www. = definite = hiatus = n	mens inc nifera found hiatus uncertain	let = specifical poor prese effects and	ly indeterminate b rvation due to sol 3/or diagenesis.	ecause of ution
TABLE 3	: DISTRIE (with en	UTION of BENT vironmental asse	HONIC FORAMINIF	ERA in Volador - :	l son: Fefer Tab	le 4).	

David Taylor, April 12, 1983.

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<u>ا</u>			GRO	SS FO	RAM	INIF	ERAL		MINOR GRAIN COMPONENTS					ITS	GROSS LITHO		1PALEO-	Ι				
7		A	ISSE	MBLAC	<u> </u>	HARA	CIER		lexc1	OCEN ud	IC forams	IN	ORGAN	IC	-GRAIN COMPONENT	rs	ENVIRONMEN ASSESSMENT	TAL				
	CORES Retres	count	s benthonic forams	s benthonic forams	c forams	RVATION Quality	RVATION - Factor	REGIME	sc t spheres rite	Dts	gments ines ules		z. ts	nodules	π: calc si marl é biomicr V: recryst biomicri calc.sil P: pyrite limonite .: f qtz s VΔ: m-c pit	lt, ite all. te f tst. f idst ited	(constant (<10m) (<0m) (<0m) (<0m) (<0m) (<200m) (<200m) (=25	RISE (=1000m) IN (2000m-1000m)	CTER CHANGE	FBI	PLANK ORAMII OSTRA	TONIC VIFERAL FIGRAPHY
	SIDEWALL Depth in	Total foram	Arenaceou	Calcareou	• planktoni	CACOJ PRESE	CaC03 PRESE	ENERGY 6 02	7 Diatom-di Biogenic py	Fish fragme Ostracods	Bryozoa fra Echinoid sp Sconce spic	Mica	c-m ang. qt glauc pelle	7 siderite	<pre>6 fracture 00: m-c ang subrd q: clayst. siltst. 0% 50%</pre>	d qtz tz. 5 100%	ESTUARINE/L INNER SHELF MID SHELF (OUTER SHELF SHELF EDGE UPPER SLOPE UPPER SLOPE UPPER SLOPE	CONTINENTAL OCTANIC BAS	E-LOG CHARA	ZONE	Depth at Base	AGE
	1325.0. 1475.0.	2000 ?	<1 7	5 7	95 7	EX P	REX	SORT		,	. .						A A			A-4 B-2	-1325 -1475	PLIOCENE LATE MIOCENE
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