

PALYNOLOGICAL ANALYSIS

OPAH-1, GIPPSLAND BASIN

by

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SUMMARY

Twenty-one sidewall cores were processed for palynomorphs in Opah-1. Examination of the samples gave the subdivision of the Latrobe Group summarised on the following table :

Unit	Zone	Depth (in feet)	Age
Lakes Entrance Formation	P. tuberculatus	7870 to 7893	Early Oligocene
	UNCONFORMITY	at 7898	
Undated Unit ?Gurnard Formation	Barren Zone	7895 to 7925	?
Flounder Formation Equivalent	P. asperopolus	7930 to 7935	Early Eocene
Latrobe Coarse Clastics	Upper L. <u>balmei</u> (W. <u>homomorpha</u> dinoflagellate Zone	7940 to 8150	Late Paleocene

TD : 8210

DISCUSSION OF ZONES

All samples processed and examined are listed on Table 1. The species identified from the samples are given on the accompanying distribution charts. The basis for choosing the zone boundaries is discussed in the following :

Proteacidites tuberculatus Zone 7870 to 7893 feet.

The consistent presence of the spore <u>Cyatheacidites</u> annulatus in the five samples clearly refers this interval to the <u>P</u>. <u>tuberculatus</u> Zone. A number of undescribed dinoflagellates in the samples also support a <u>P</u>. <u>tuberculatus</u> Zone age which would correlate the section with the basal <u>Lakes Entrance Formation</u> in other wells. The grey calcareous siltstone lithology of the interval, and the E-log response would also indicate correlation with the Lakes Entrance Formation.

In contrast to this palynological data, the results from the study of foraminifera (Taylor 1977b) suggests that the interval between 7850 to 7893 feet contains both J2 and K zone faunas, which would make the section partially Late Eocene in age and also time equivalent to the youngest part of the Gurnard Formation in some other wells. This is best demonstrated by the data obtained from the recent well Swordfish-1 (Taylor 1977a, Partridge 1977). In this well, Early Oligocene (zone J2) and Late Eocene (zone K) faunas were identified over an interval of 45 feet at the top of the Gurnard Formation (in sidewall cores between 6560-6564 feet and 6571-6604 feet respectively). However, on palynology these same samples were referred respectively to the Upper <u>N. asperus</u> Zone and Middle N. asperus Zone. The presence of Late Eocene in this section in Opah-1 can be disputed as it is based on a doubtful identification of Subbotina linaperta in the sidewall core at 7890 feet. The presence of the Early Oligocene J2 zone cannot be disputed nor the correlation with Swordfish-1 as the same planktonic foraminiferal assemblage occurs in both wells. The conflict with the palynological data can be resolved if we consider that the base of the P. tuberculatus Zones lies within the J2 planktonic foraminiferal zone and not at the base of J1 as has been previously stated (Partridge 1976, figure 2). The implication of this to the geology of the Gippsland Basin is that fine grained Lakes Entrance Formation sedimentation commences slightly earlier than previously recognised (Partridge 1976, figure 7) and approximates even closer the Eocene-Oligocene boundary.

Barren Zone (sidewall cores from 7895 to 7925 feet).

The five sidewall cores processed in this interval yielded mainly flaky brown amorphous organic material and some black angular woody material but no palynomorphs. Lithologically, this section of brown to rare green silty sandstone with accessory glauconite and limonite is similar to the Gurnard Formation. It is unusual in being oxidised, which is the most likely reason for the absence of spore-pollen and dinoflagellates.

The section is reminiscent of the 130 feet of red to brown ferruginous shales and siltstones in Moray-1, between 5370 and 5500 feet (-5338 to -5470 feet subsea). Like Opah-1, the Moray-1 section lacks firm age dating. In Moray-1, the foraminiferal zone I-1 is identified down to 5360 feet. No foraminifera were identified in sidewall cores below this and new fauna observed in cuttings gave only an Eocene to Early Oligocene age. The section was also barren of palynomorphs except for a single sidewall core at 5490 feet which gave a Lower <u>N. asperus</u> age. In the Moray-1 well completion report, most of this section was correlated with the early Oligocene. It could just as likely be Middle to Late Eocene in age and correlated with the Gurnard Formation which is also the obvious correlation for this section in Opah-1.

Proteacidites asperopolus Zone 7930 to 7935 feet.

The two sidewall cores are assigned to the <u>P. asperopolus</u> Zone on the presence of <u>Myrtaceidites tenuis</u>, <u>Conbaculites apiculatus</u> and <u>Santalumidites</u> <u>cainozoicus</u>. The dinoflagellates identified support this age, especially the occurrence of <u>Homotryblium tasmaniensis</u>. Unfortunately, only a very low yield of fossils was recovered from the samples. The low yield particularly hampered the identification of the dinoflagellates present and consequently, the samples cannot be referred to any of the Early Eocene <u>Wetzeliclla</u> zones nor closely compared with dinoflagellate assemblages from the age equivalent Flounder Formation.

Upper Lygistepollenites balmei Zone 7940 to 8150 feet (and Wetzeliella homomorpha Dinoflagellate Zone).

Less than 200 feet of Upper L. <u>balmei</u> Zone was intersected. The nine samples examined from this interval are assigned to the zone on the common occurrence of Lygistepollenites balmei and the presence of other index species such as <u>Australopollis</u> obscurus, <u>Amosopollis</u> <u>dilwynensis</u>, <u>Camarozonosporites</u> <u>bullatus</u>, <u>Gambierina</u> <u>rudata</u> <u>and</u> <u>Polycolpites</u> <u>langstonii</u>. All these species become extinct near the top of the L. <u>balmei</u> Zone. Indicator species for Upper subdivision of the L. <u>balmei</u> Zone present in the samples include <u>Banksieaeidites</u> <u>elongatus</u>, <u>Cyathidites</u> <u>gigantis</u> (fairly common), <u>Malvacepollis</u> <u>diversus</u> and <u>Verrucosisporites</u> <u>kopukuensis</u>. Two unusual occurrences are <u>Bysmapollis</u> <u>emarciatus</u> and <u>Gephrapollenites</u> <u>cranwellae</u>. Their presence in this interval extends the lower limits of their ranges from the Lower M. <u>diversus</u> Zone to the Upper L. <u>balmei</u> Zone.

The dinoflagellates present support the Upper L. balmei age. The presence of Wetzeliella homomorpha refers the section to the dinoflagellate zone bearing that name. The 54 feet of siltstone between 7939 and 7993 feet from the upper part of this zone is anomalous and was unexpected compared with the sandy facies in the Upper L. balmei Zone in surrounding wells. The six sidewall cores examined from this unit all yielded good zone assemblages and are closely comparable with assemblages from the underlying section (7993-8210 feet), and from adjacent wells. The possibility that this upper unit could be stratigraphically younger and contain reworked L. balmei Zone assemblages is emphatically rejected based on the absence of the common marker species of the M. diversus or P. asperopolus Zones. The occurrences of Bysmapollis emarciatus and Gephrapollenites cranwellae in Opah-1 cannot be taken as evidence of reworking as both forms are very minor components of assemblages from younger zones. If they were reworked, the Opah-1 assemblages should contain the commoner M. diversus and P. asperopolus Zone indicator species as well.

Taylor (1977b) does however, record the planktonic foraminifera <u>Subbotina</u> <u>frontosa</u> from the sidewall core at 7970 feet within this upper section. The worldwide range of this species suggests that it does not extend below the Early Eocene. Further, the maximum possible age for the section based on this single planktonic species conflicts with the palynology which indicates a Late Paleocene age. This age for the palynological zones is derived by correlation with dinoflagellate assemblages in New Zealand, again dated by planktonic foraminifera. Considering that both the spore-pollen and dinoflagellate assemblages in this section in Opah-1 are represented by a number of species the age dating and correlation based on the palynology is considered more reliable than that of the single planktonic foraminifera species.

REFERENCES

Partridge, A.D., 1976, The Geological expression of Eustacy in the Early Tertiary of the Gippsland Basin: APEA J., v.16, pt.1, p.73-79.

Partridge, A.D., 1977, Palynological analysis Swordfish-1, Gippsland Basin : Esso Aust. Ltd. Palaeo. Rept., 1977/13.

Taylor, D., 1977a, Foraminiferal sequences Swordfish-1 : <u>Esso Aust.</u> Ltd. Palaeo. Rept., 1977/5.

Taylor, D., 1977b, Foraminiferal sequence Opah-1 : Esso Aust. Ltd. Palaeo. Rept., 1977/12.

TABLE - 1 : SUMMARY OF PALYNOLOGICAL ANALYSIS, OPAH-1, GIPPSIAND BASIN

.

SAMPLE AND DEPTH	ZONE	AGE	CONFIDENCE RATING	YIELD	DIVERSITY	REMARKS
SWC 61 7870'	P. tuberculatus	Early Oligocene	0	Moderate	Moderate	Cyatheacidites annulatus present
SWC 30 7880'	11	11 11	0	Moderate	High	11
SWC 29 7885'	u	11 11	0	Moderate	Moderate	π
SWC 28 7890'	"	87 89	0	Moderate	Moderate	n
SWC 27 7893'	"	11 · 11	0	Low	Moderate	11
SWC 26 7895'	Barren of Fossil	S				Only brown amorphous material of rounded shape recovered.
SWC 23 7902'	Barren	•				Nothing recovered
SWC 22 7905'	Barren of Fossil:	5				Brown flaky material of indeterminar origin recovered.
STAC 18 7920'	Barren of Fossils	S				Angular woody material recovered
SWC 17 7925'	Barren of Fossil	S				Brown flaky and woody materials recovered.
SWC 16 7930'	P. asperopolus	Early Eccene	1	Low	Moderate	
SWC 15 7935'	P. asperopolus	Early Eccene	1	Very Low	Moderate	
SVC 14 7940'	Upper L.balmei	Late Paleocene	0 ·	High 🥳	Moderate	W. homomorpha Dino. Zone
SWC 13 7945'	TT TT	11 II	0	High	High	"
SWC 12 7950'	11 11	11 11	0	High	High	11
SWC 10 7970'	11 11	" "	0	High	Moderate	n
STAC 9 7980'	11 11	** **	0	High	High	17
SWC 8 7992'	11 11	11 11	0	High	High	11
SWC 5 8070'	11 11	11 17	0	Moderate	High	Π
SWC 3 8106'	18 18	11 11	0	High	High	
SWC 1 8150'	FF - FF	11 11	0	High	High	

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BASIN	GIPP	SLAND			DAT	E	JUNE	_20,	1977		
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AGE	PALYNOLOGIC ZONES	Preferred Depth	Rtg.	Alternate Depth		2 way time	Preferred Depth	Rtg	Alternate Depth		2 way time
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<u>0-</u>	U. <u>N. asperus</u>										
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1	L. <u>N. asperus</u>					•					
NE	<u>P</u> . <u>asperopolus</u>	7930	1				7935	1			
EOCENE	U. <u>M</u> . <u>diversus</u>						•	•			
	M. <u>M</u> . <u>diversus</u>										
· 	L. <u>M</u> . <u>diversus</u>					2, 2					
ENE	U. <u>L</u> . <u>balmei</u>	7940	0				8150	_1			
PALE OCENE	L. <u>L. balmei</u>						·				
VI III	<u>T. longus</u>										
ر م	<u>T. lilliei</u>	·•									
CEOU	<u>N. senectus</u>										
LATE CRETACEOUS	<u>C. trip./T.pach</u>										
υ	<u>C. distocarin</u> .										
EA	<u>T. pannosus</u> RLY CRETACEOUS										
	ALI CREIACEOUS										
PR	E-CRETACEOUS										
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COMM	ENTS:				Dino.	Zone 79	940-8150 fe	et (1	rating 1)		
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	pollen	or microp	lankto	n.			n zone spec				
	and/or	microplank	ton.			•	n non-diagn				
	pollen	or micropl	lankto	n, or both	•		one species				
		GS, <u>NO CONI</u> Lankton.	<u>IDENC</u>	<u>E</u> , assembla	age w	ith non-	-diagnostic	spoi	res, poller	n and	/or
NOTE	: If a sample ca Also, if an er better confide	try is giv	ven a	3 or 4 con	fiden	ce ratin	ng, an alte	enti rnate	cy should l e depth wit	be ma th a	de.
DATA	A RECORDED BY:	A.D. Par	<u>tridg</u>	<u>e</u>			June,				-
DATA	REVISED BY:					DATE					-

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*C=core; S=sidewall core; T=cuttings.

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 $C = core_i$ S = sidewall core_i T = cuttings.

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*C=core; S=sidewall core; T=cuttings.

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▼C≈core; S≈aidewall core; T ≈ cuttings.